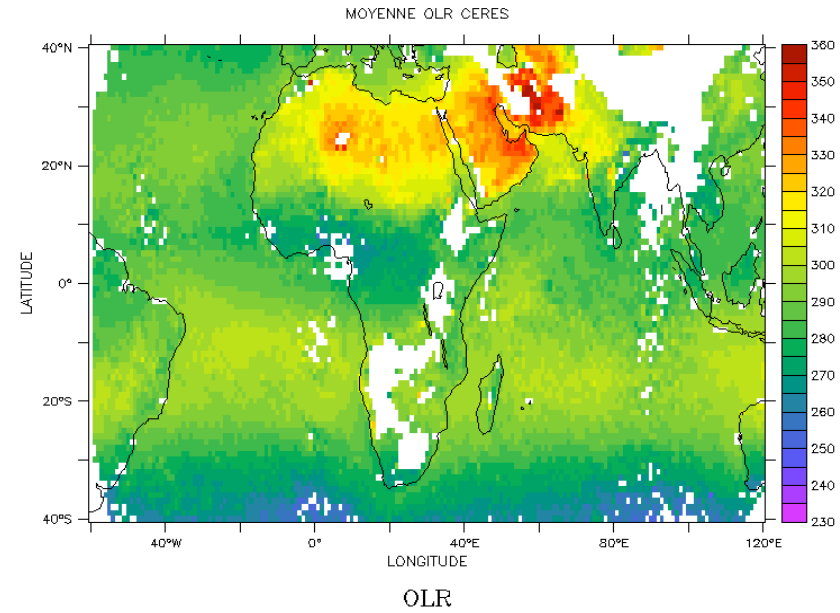
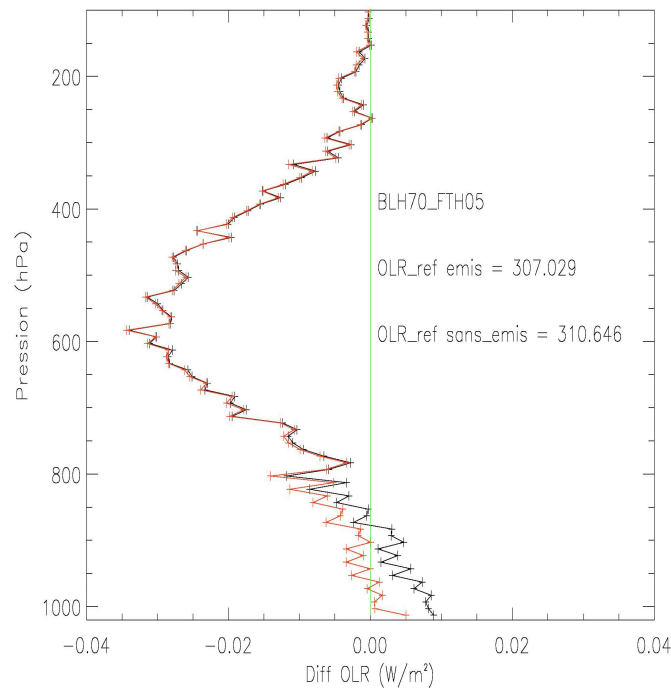


# Key parameters to estimate the clear-sky OLR over tropical regions: Simple models and their evaluation

R. Guzman, L. Picon, R. Roca

LMD, CEET team, N. Gif, O. Chomette, P. Raberanto



# Outline

## I – Introduction

I.1 State of the art

I.2 Motivation

## II – Radiative transfer tool : Modtran

II.1 – Idealized atmospheres, starting assumptions

II.2 – Surface emissivity impacts over the IR radiation

## III – Statistical approach : Satellite data

III.1 – Available data sets

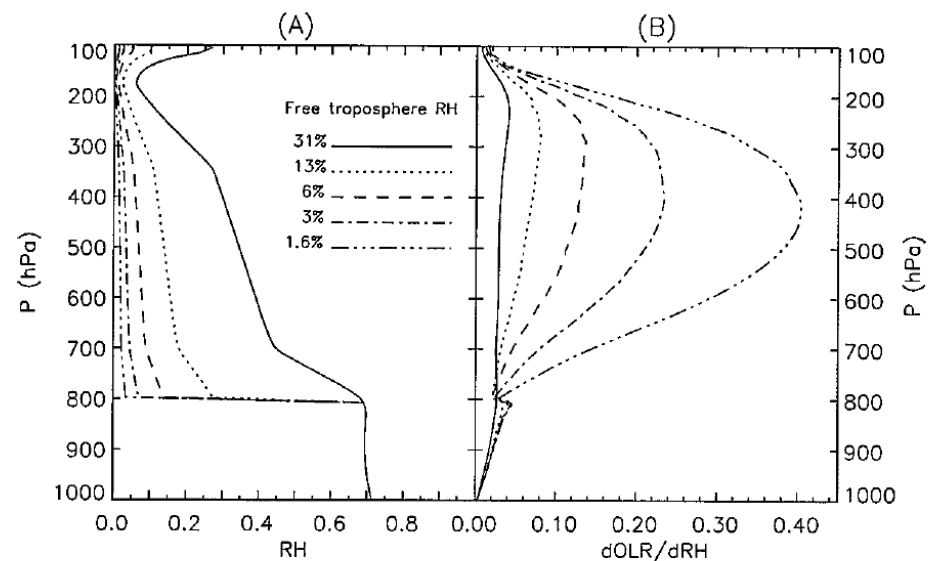
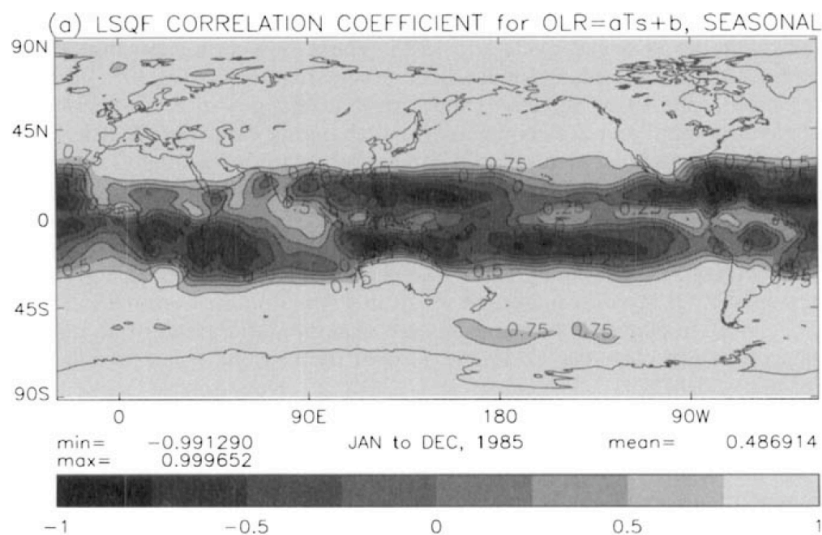
III.2 – Multi-linear regression

## IV – Conclusions et perspectives

# Introduction (1/3)

## I.1 – State of the art

- Infrared (IR) Outgoing Longwave Radiation (OLR) is strongly correlated to Surface Temperature ( $T_s$ ) over the whole globe (*Allan et al. 1999*).
- In the tropical regions, moist plays a major role on OLR as well, particularly on its variability (*Buehler et al. 2004*). OLR is mainly sensitive to the mid-upper tropospheric humidity (*Spencer et Braswell 1997, Allan et al. 1999*).



# Introduction (2/3)

## I.1 – State of the art

- In the tropics there are surfaces with low emissivities in the IR ( $\sim 0.85$ ), mainly the deserts, and that are not accurately estimated (*Zhang et al. 2007*).
- Evaluate if the blackbody emissivity approximation can yield systematic errors that should be taken into account for the IR radiation estimates.

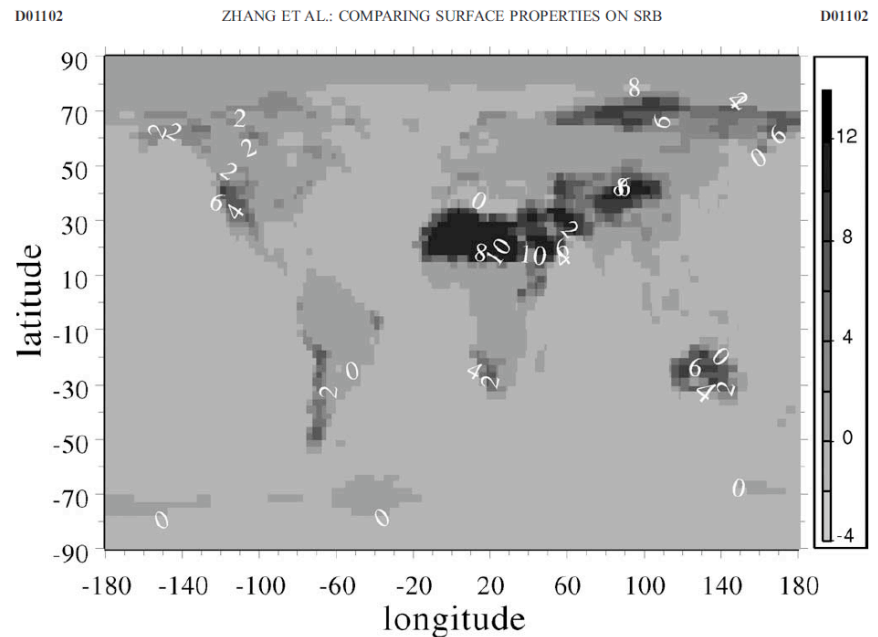
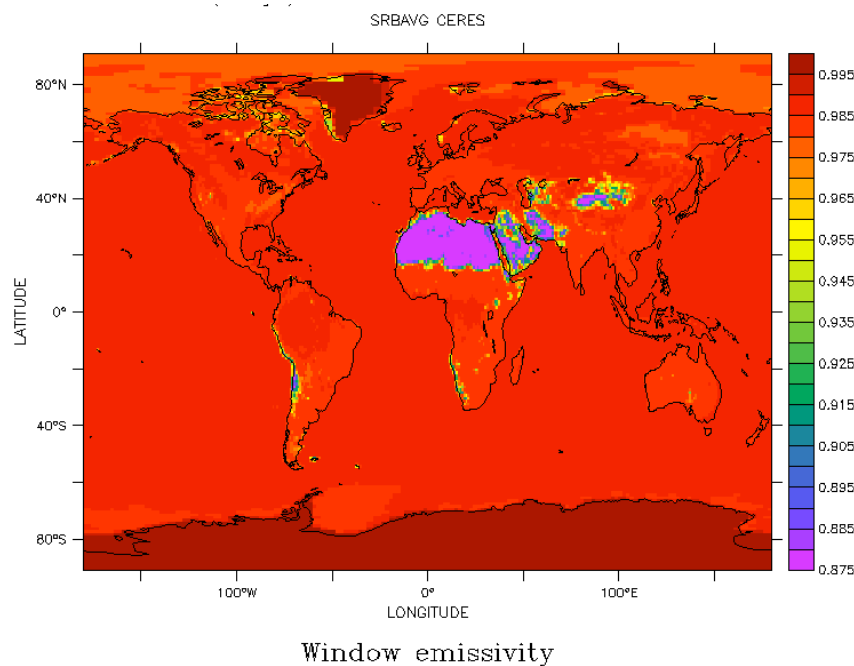


Figure 7. Global map of differences (ISCCP-FD minus SRB) of window region emissivities (in %) from ISCCP-FD 1992 annual mean and SRB climatology.

# Introduction (3/3)

## I.2 – Motivation

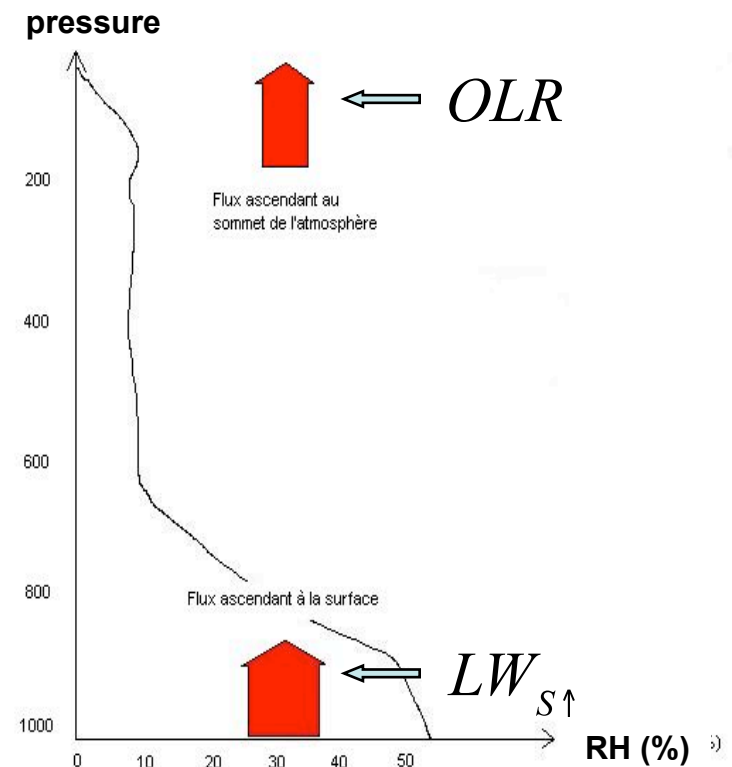
### Clear-sky greenhouse effect

- To understand and estimate the clear-sky greenhouse effect over the tropics
- Create a simple model of  $G_a$  to evaluate the impact of different kind of changes on the key variables of  $G_a$  or  $OLR$ .

Two complementary ways to achieve that :

- Idealized study cases with the Modtran model
- Statistical approach with satellite data

$$G_a = LW_{S\uparrow} - OLR$$



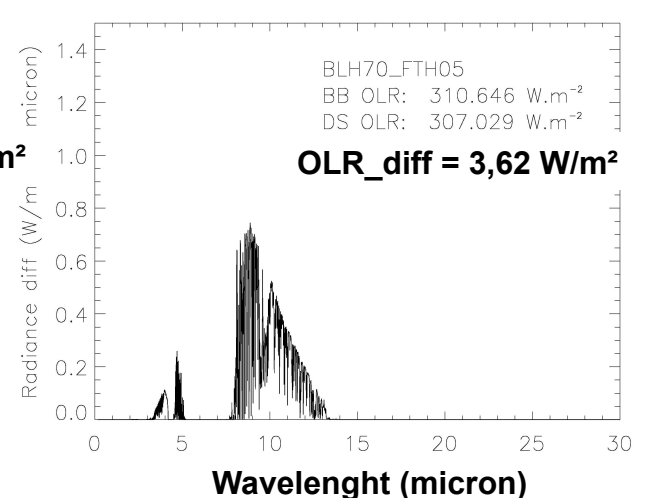
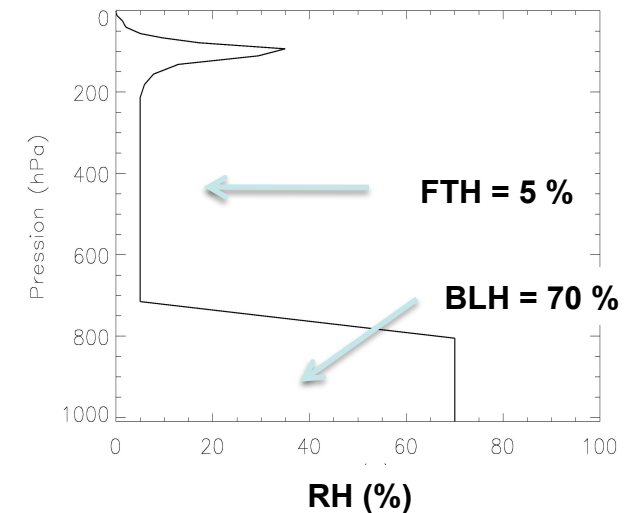
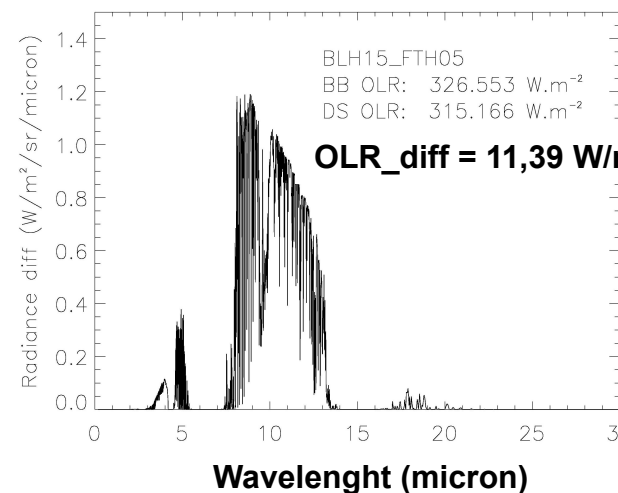
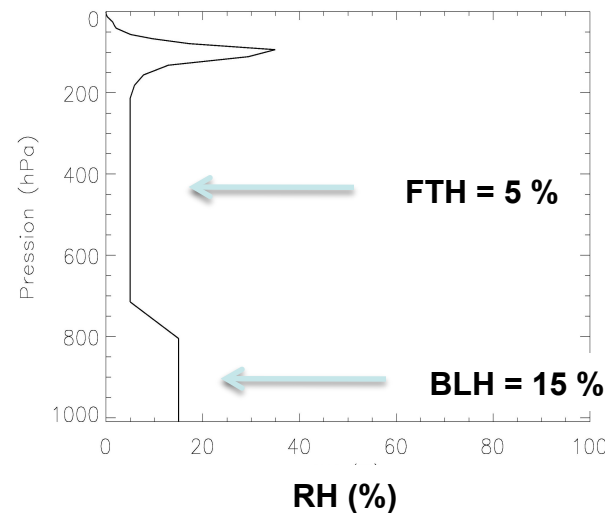
# Modtran (1/3)

## II.1 – Idealized atmospheres, starting assumptions

- Two humidity layers : the Boundary Layer (BL) and the Free Troposphere (FT) layer.

- All other atmospheric parameters are from the McClatchey tropical standard profile (*McClatchey et al. 1971*)

- There is a 3% difference in the OLR (11 W/m<sup>2</sup>) for the driest atmosphere.



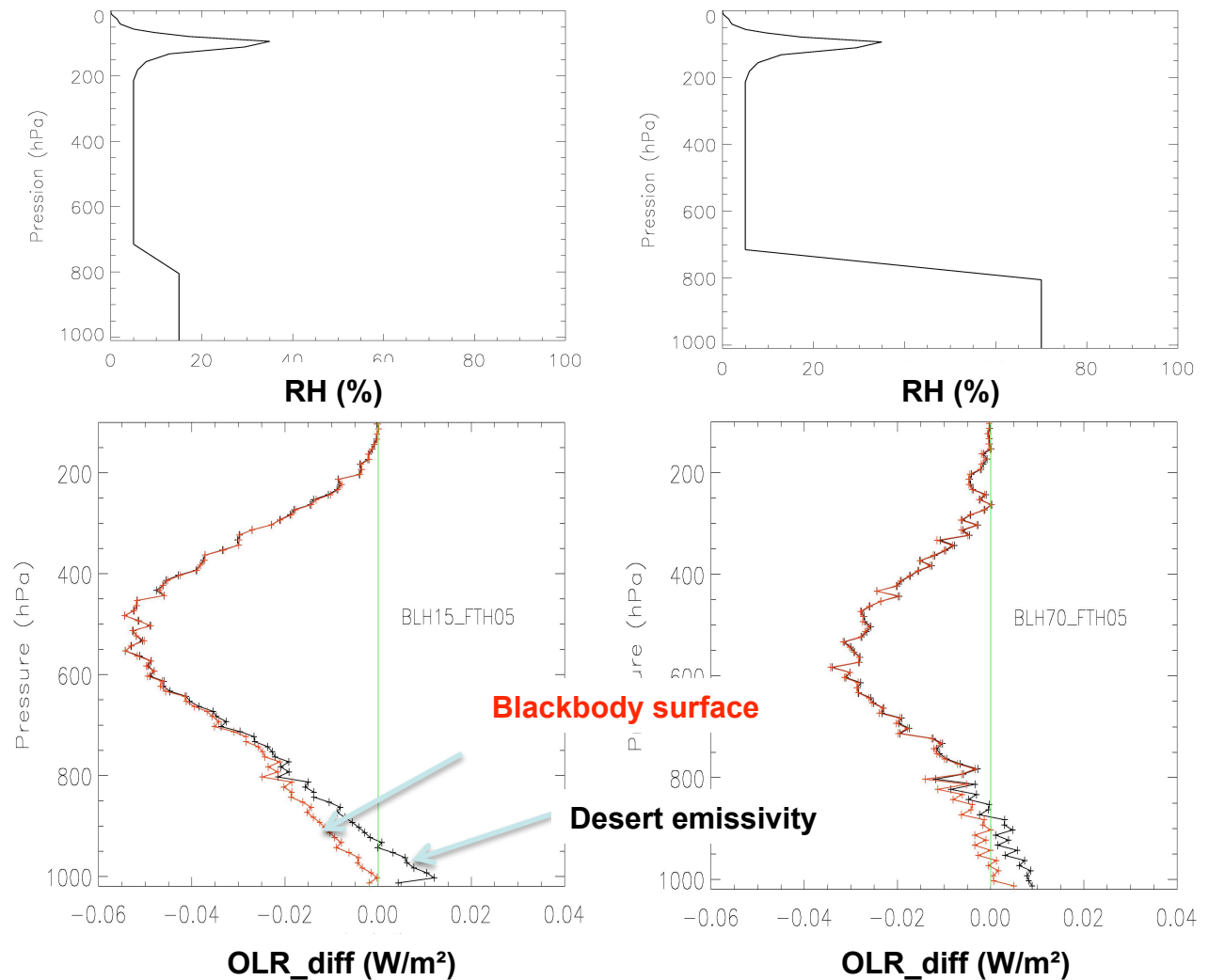
# Modtran (2/3)

## II.2 – Surface emissivity impacts over the IR radiation

OLR jacobians with respect to RH for the two types of surfaces

+1% perturbations in RH over 10 hPa thick layers

Differences between the two surfaces appear in the lower part of the atmosphere



# Modtran (3/3)

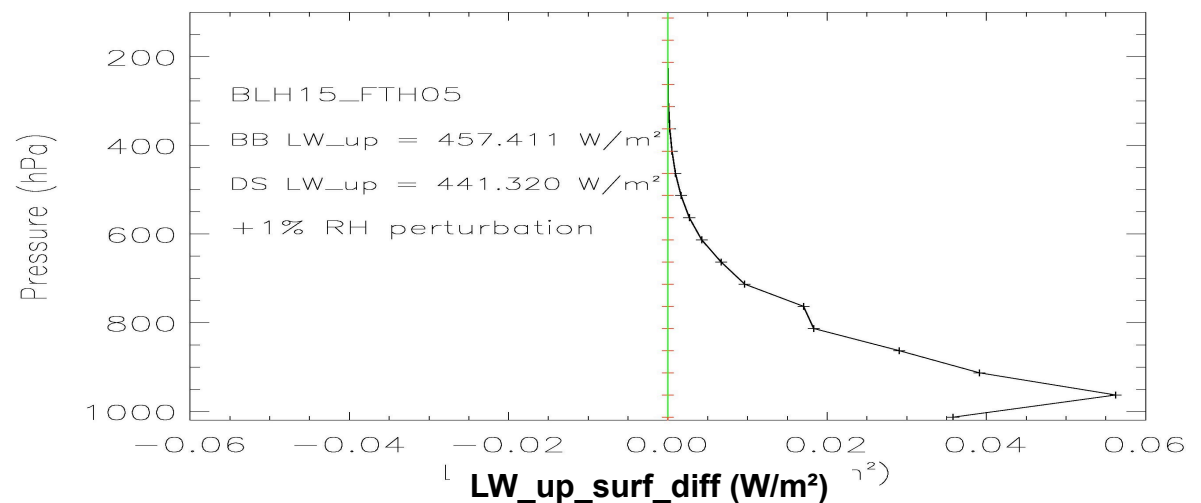
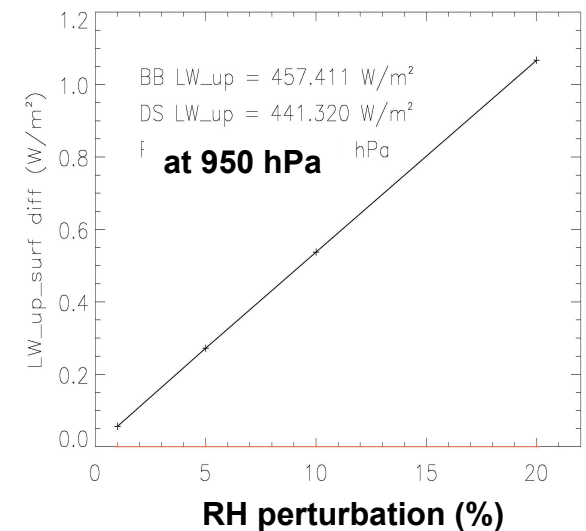
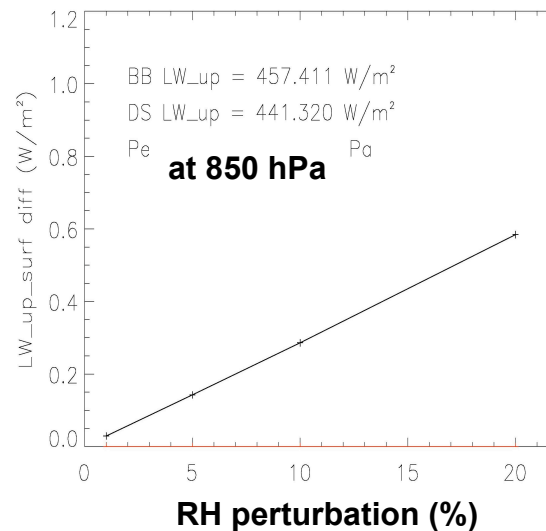
## II.2 – Surface emissivity impacts over the IR radiation

For desert-like surfaces :

- Linear increase of the surface upward IR flux for greater RH perturbations

- The slope gets bigger if the perturbed layer gets closer to the surface

- We plot the LW\_up surface flux jacobian with respect to RH perturbations





# Satellite data (1/6)

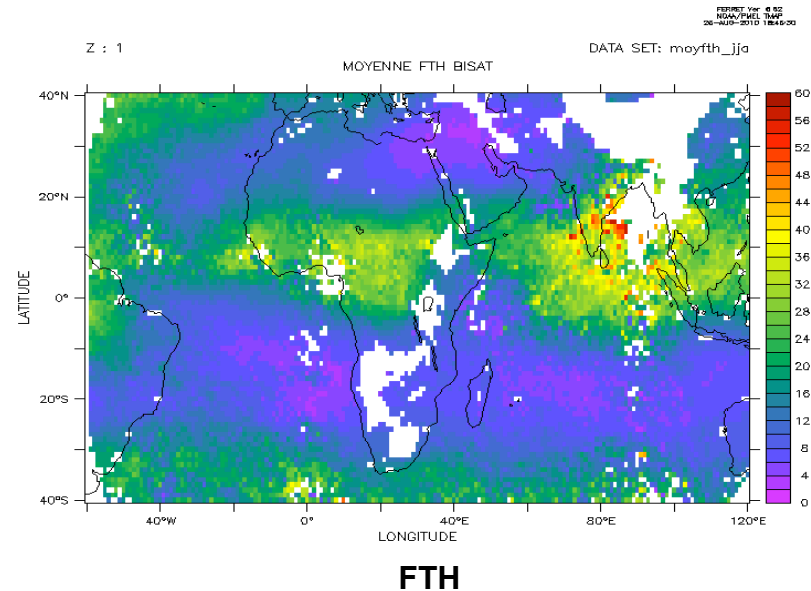
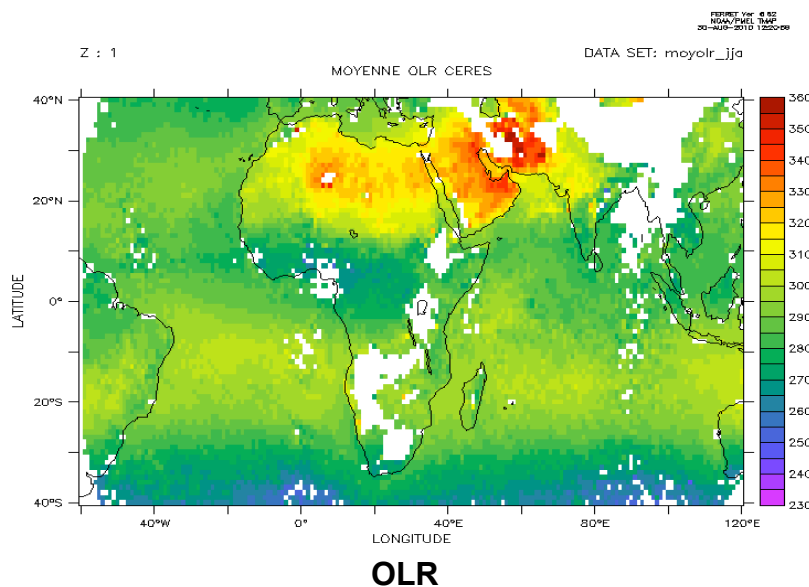
## III.1 – Available data sets

- **CERES SSF** – 07/2002-06/2003 (AQUA)

We take from this data sets the following products : **OLR** and **Ts**

- **FTH BISAT** – 2000-2005 (2 x METEOSAT)

FTH estimate, every 3 hours, over the bisat region (60°W-120°E , 40°S-40°N)  
(water vapour band product, METEOSAT (*Brogniez et al.2006*))

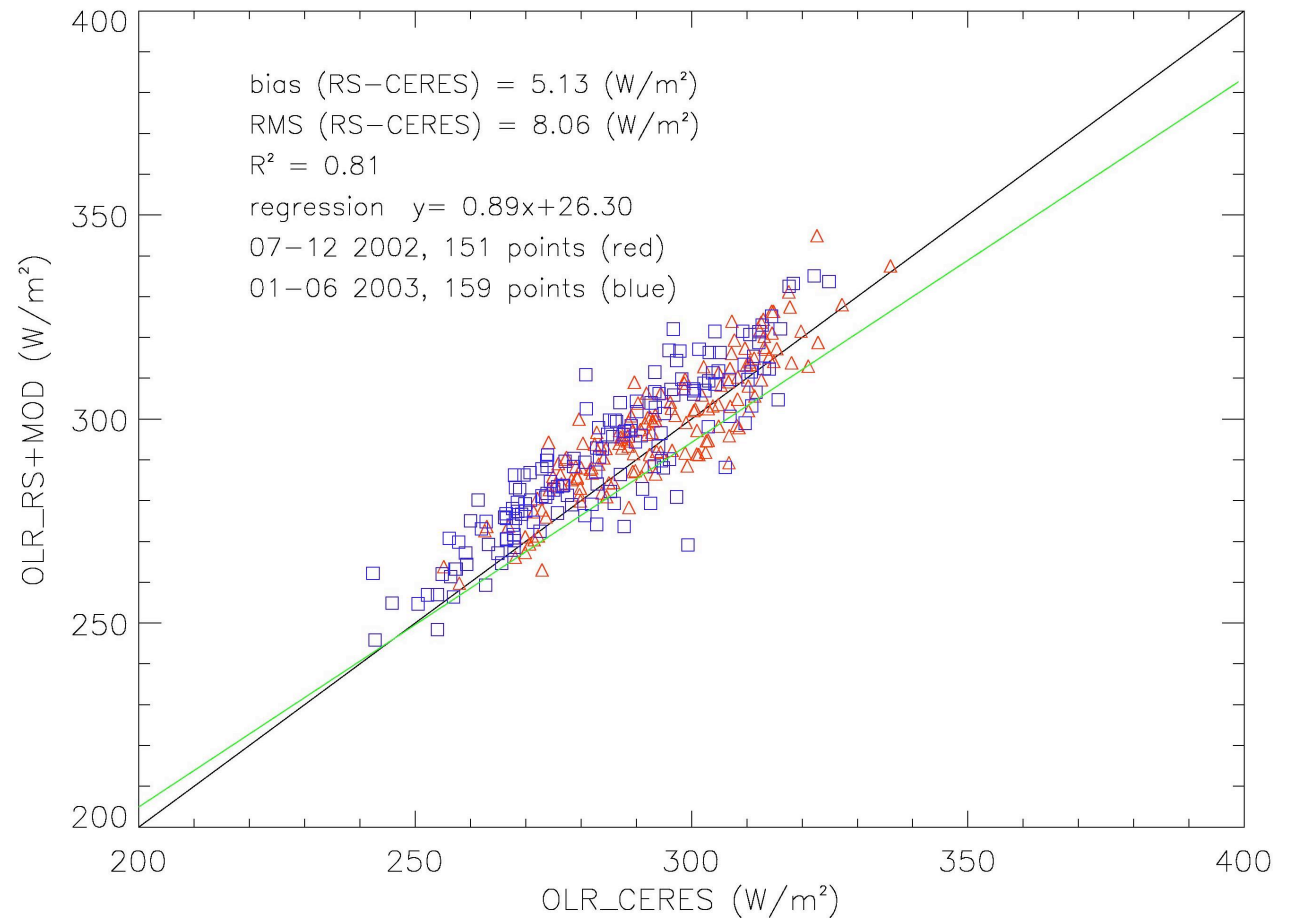


# Satellite data (2/6)

## III.1 – Available data sets

### Night OLR data

We plot the OLR computed by Modtran with radiosonde humidity profiles vs CERES OLR (space time coincidence :  $1^\circ \times 1^\circ$ , two hours time bin).



**The « best » OLR estimate, vertically and spectrally resolved**

# Satellite data (3/6)

## III.2 – Multi-linear regression

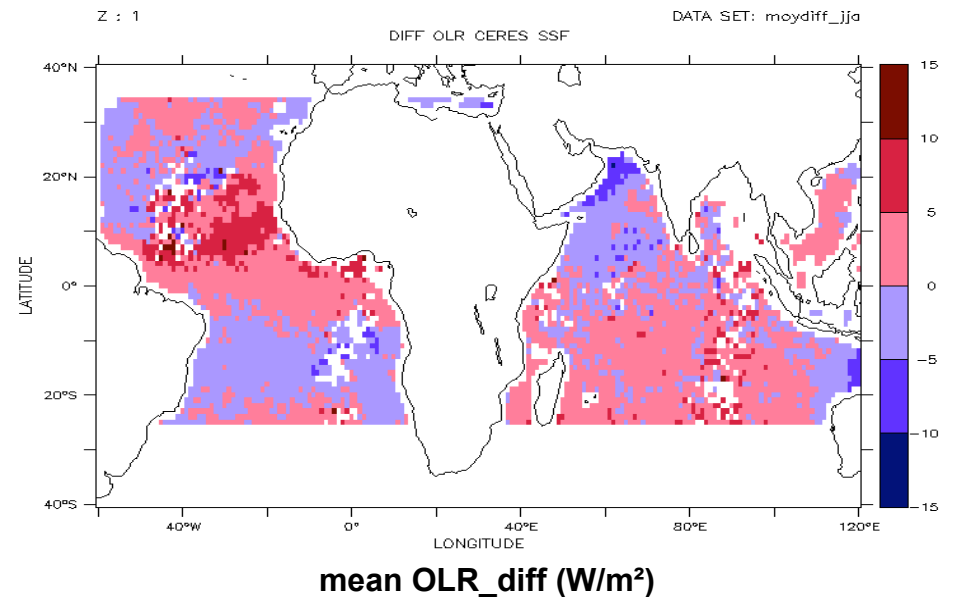
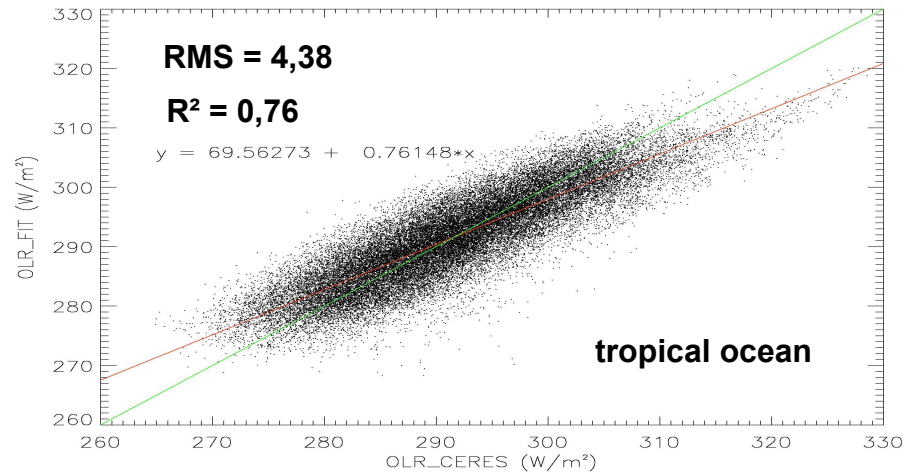
We focus on the JJA season,  
25°S-35°N and night  
measurements of OLR

We fit the OLR with the following  
model:

$$OLR = a \cdot \sigma \cdot T_s^4 + b \cdot \ln(FTH) + c$$

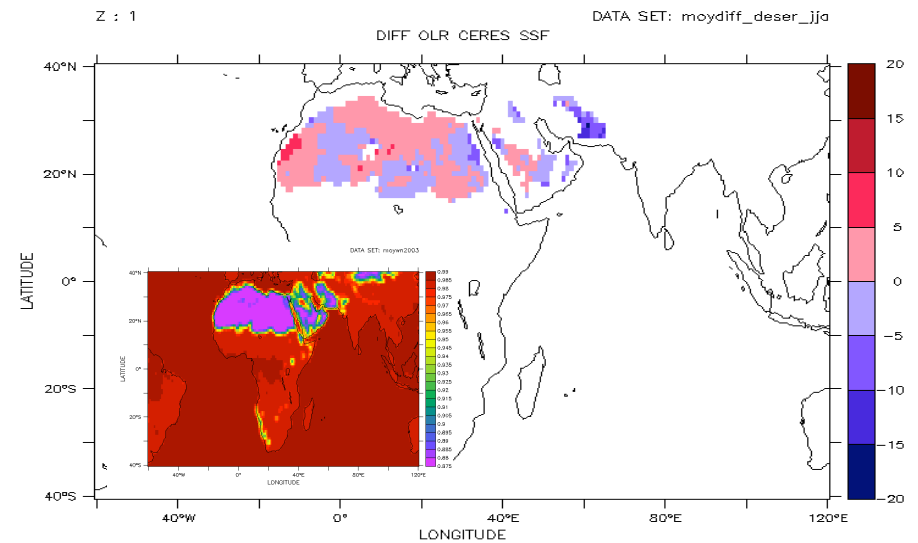
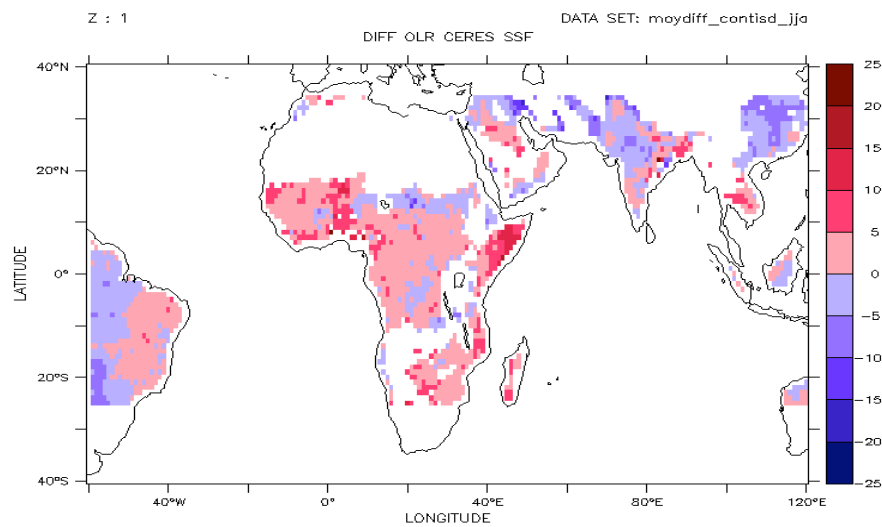
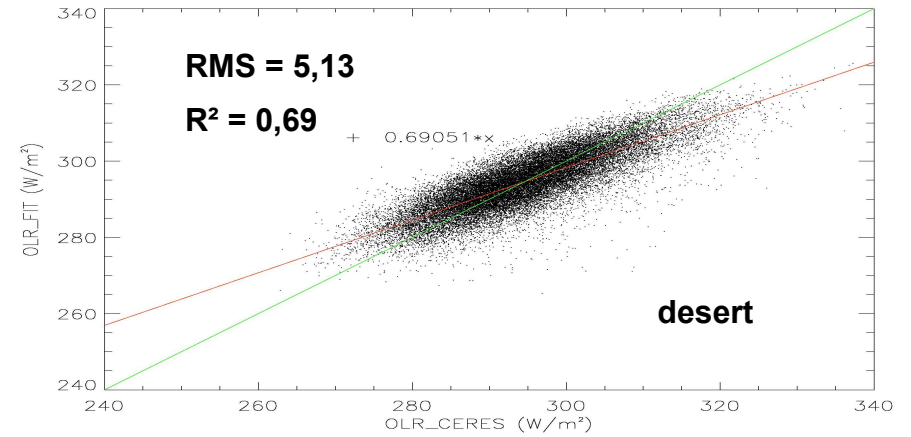
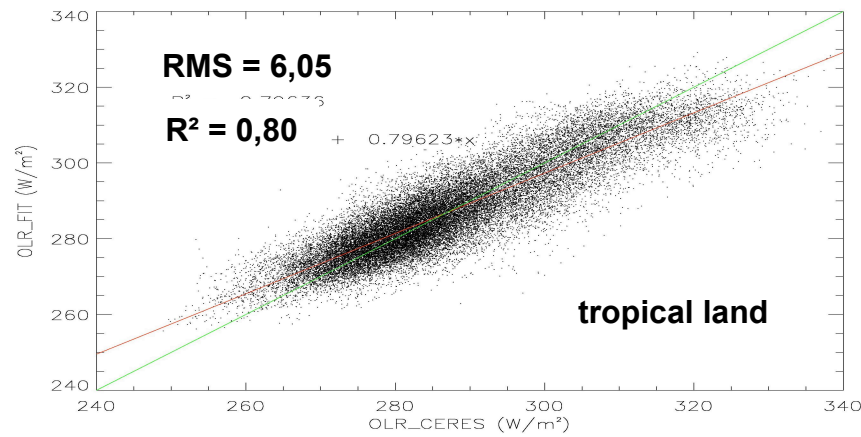
Ocean fit vs data (top)

Mean of the difference between  
the fit and the data (bottom)



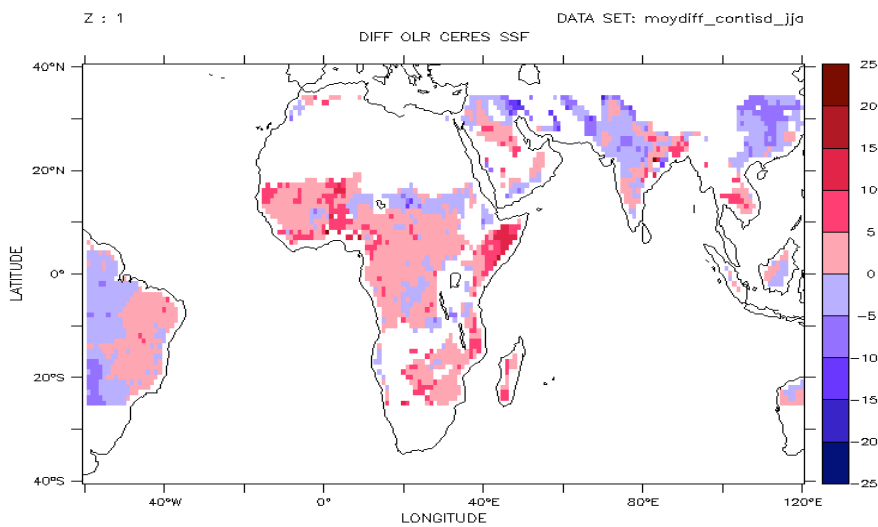
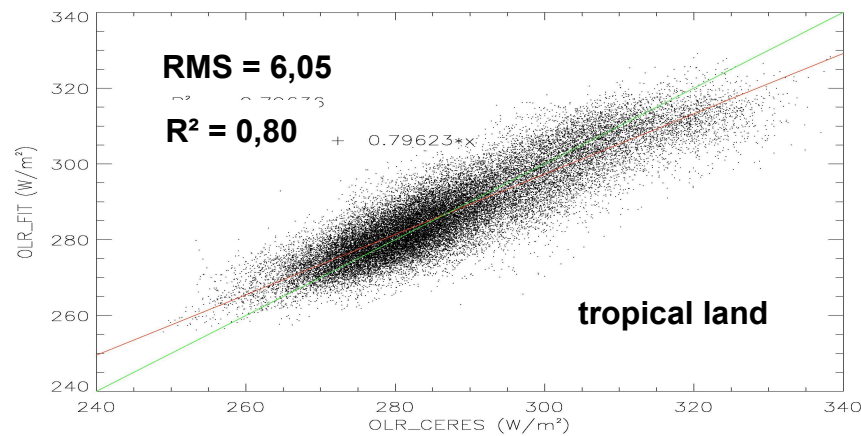
# Satellite data (4/6)

## III.2 – Multi-linear regression

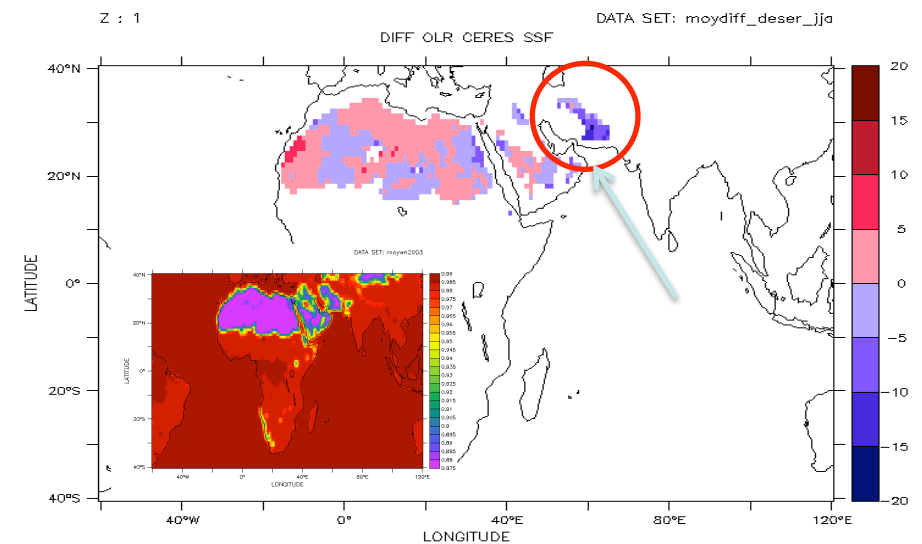
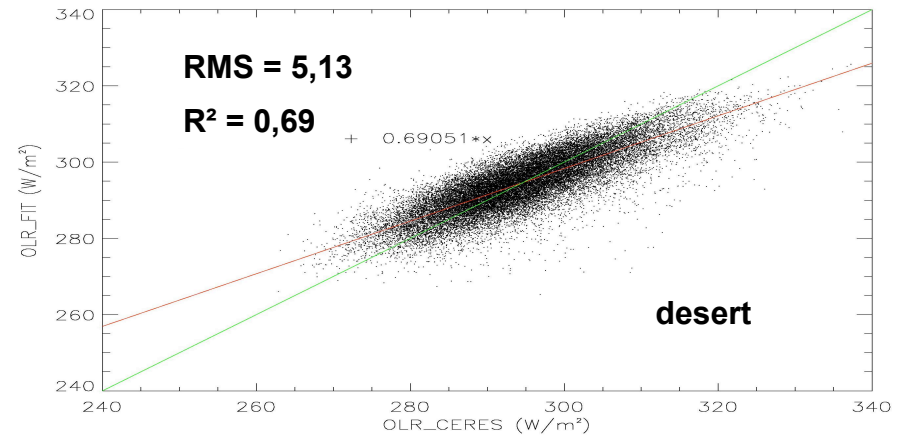


# Satellite data (4/6)

## III.2 – Multi-linear regression



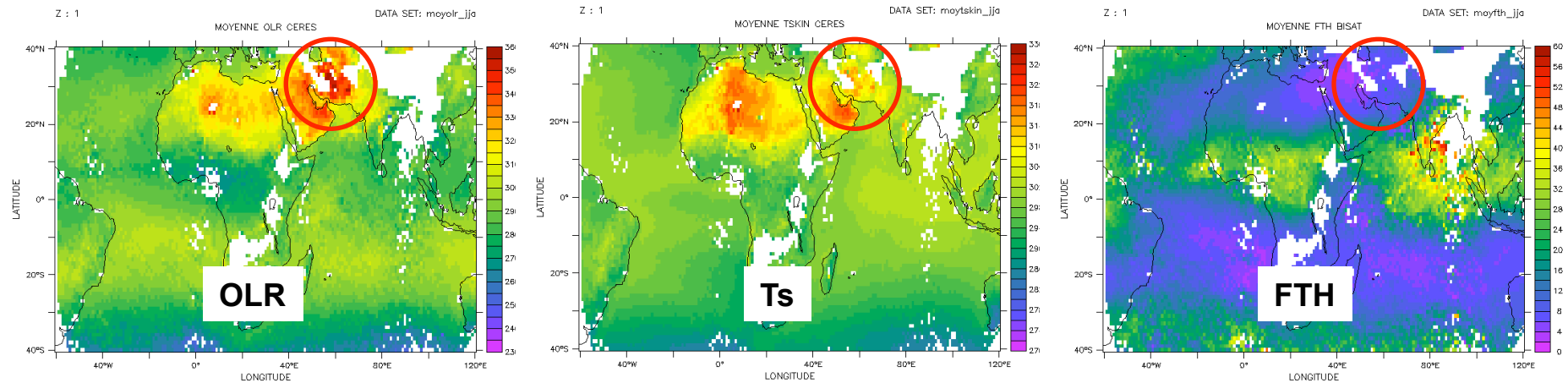
mean OLR\_diff ( $W/m^2$ )



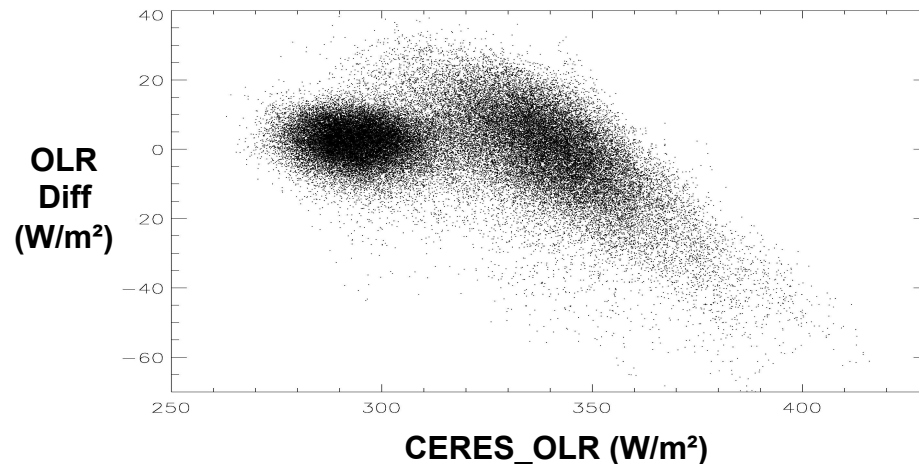
mean OLR\_diff ( $W/m^2$ )

# Satellite data (5/6)

## III.2 – Multi-linear regression



Mean of the selected data : OLR (left), Ts (centre) and FTH (right).  
The highest OLRs do not correspond to the highest temperatures



Problem for the two-parameter model to estimate the highest OLRs (right cluster, day data)

# Satellite data (6/6)

## III.2 – Multi-linear regression

Correlation coefficients between each variable and the OLR (night)

R : linear  
correlation  
coefficient  
between the FIT  
and the DATA

Stddev : standard  
deviation of the  
FIT-DATA

	Correlation coefficients				
Surface type	$\sigma \cdot T_s^4$	$\ln(FTH)$	$\ln(PWAT)$	$R^2$	<i>stddev</i>
Ocean	-0.009	-0.820	-0.540	0.765	4.378
Land	0.572	-0.643	-0.090	0.775	5.816
Desert	0.099	-0.790	0.327	0.693	5.135

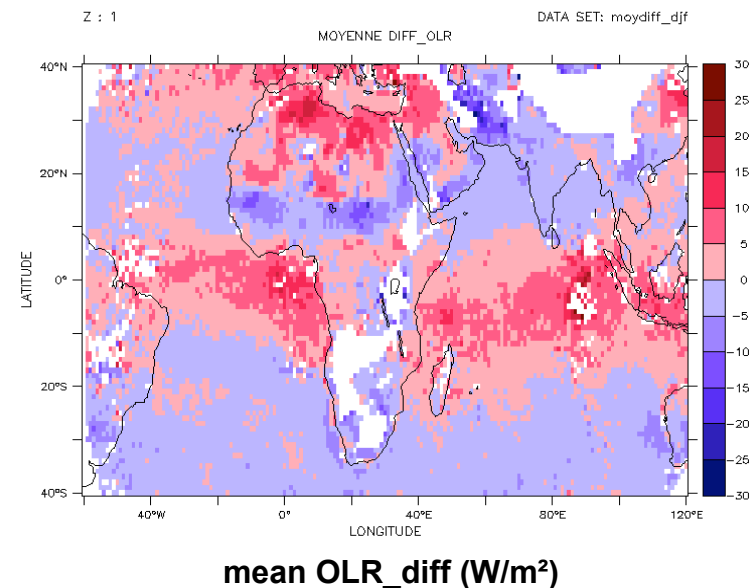
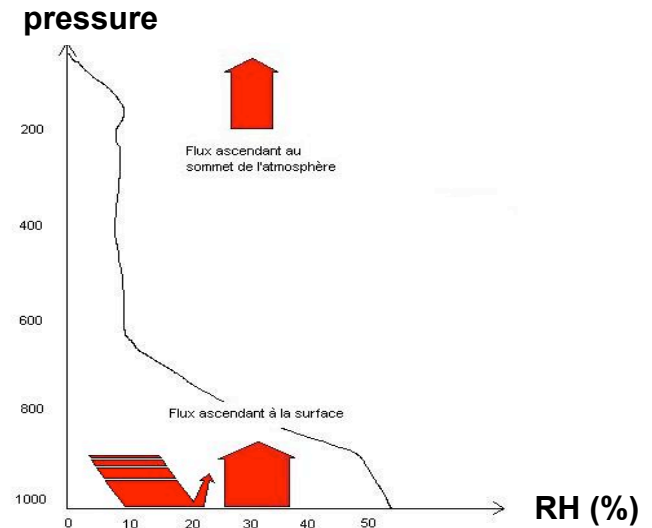
The two-parameter model reproduces the OLR with almost the same precision than the vertically and spectrally resolved model

# Conclusions

- Surface emissivity has to be taken into account when it is low, particularly in the « window » region
- Systematic errors can reach  $10 \text{ W/m}^2$  and get bigger for higher temperatures, lowest emissivities and drier boundary layers
- The two-parameter statistical model is almost as good as a « complete » vertical and spectral model

## Perspectives :

- Identify the reasons of the bias
- Improve the statistical model
- study the variability of OLR and Ga within this statistical model

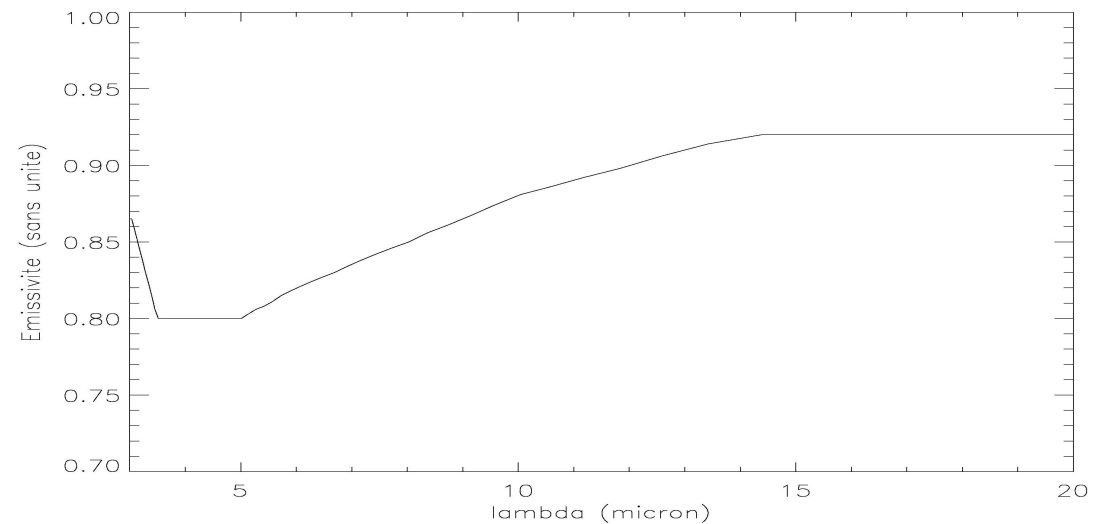




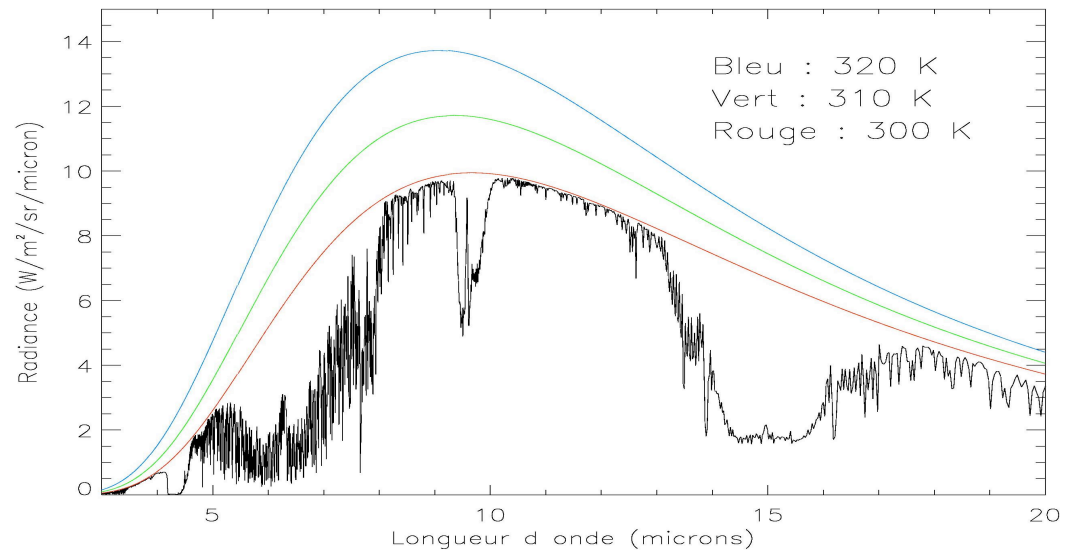
# Appendix 1

## Modtran

Desertic surface emissivity spectrum from the Modtran model. We highlight the particularly low emissivity values in the atmospheric window, between 8 and 12  $\mu\text{m}$ .



The bulk of the most intense emissions for Ts close to 300 K are in the window part of the spectrum.



# Appendix 2

## Modtran

Over desert-like emissivity surfaces :

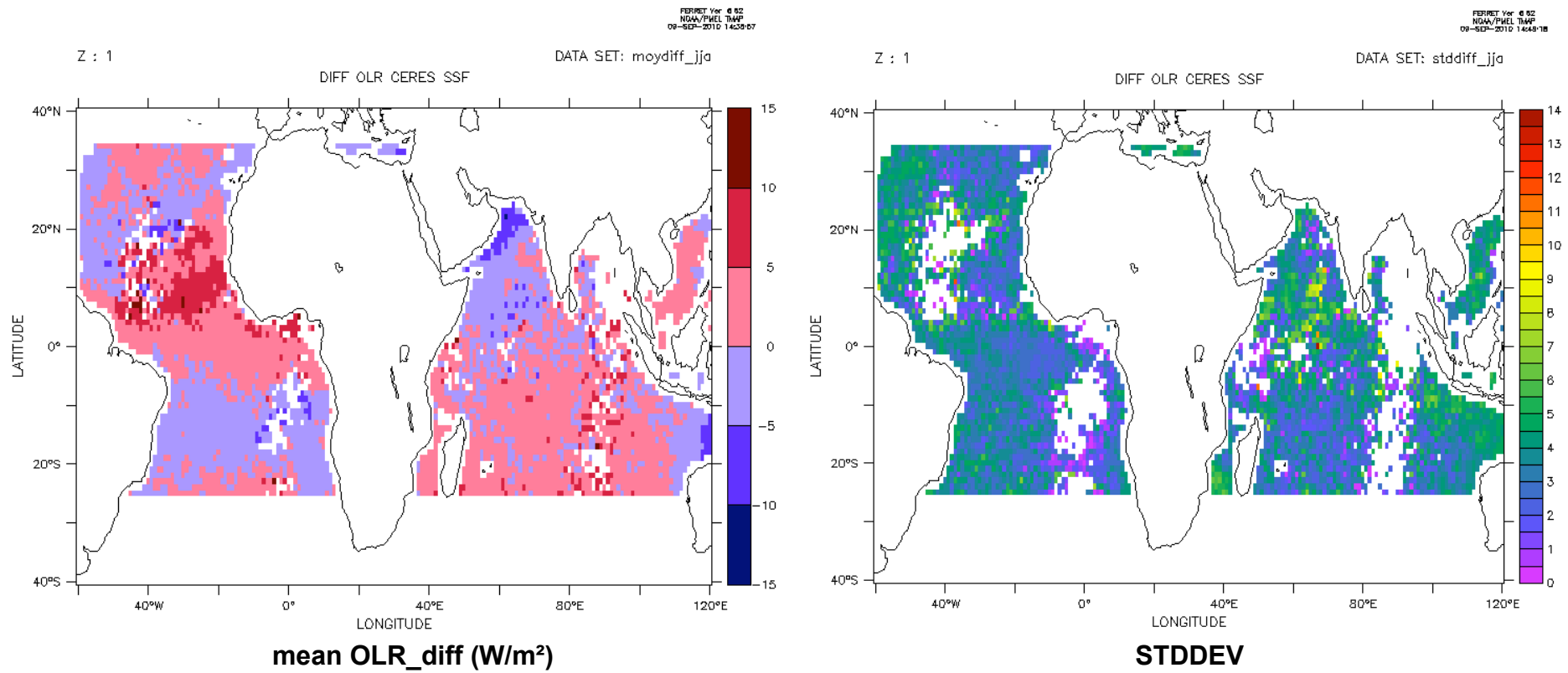
- $T_s$  variations have a major impact on the amount of emission in the atmospheric window
- OLR jacobian becomes positive for RH perturbations in the low layers of the atmosphere
- Emissivity should be taken into account for low emissivity surfaces, otherwise error of 10 W/m<sup>2</sup> or more could be done on the OLR estimate

$$LW_{S\uparrow} \approx (1 - \varepsilon) \cdot LW_{A\downarrow}(BLH, T_A) + \varepsilon \cdot \sigma \cdot T_S^4$$

$$OLR \approx f(LW_{S\uparrow}, FTH)$$

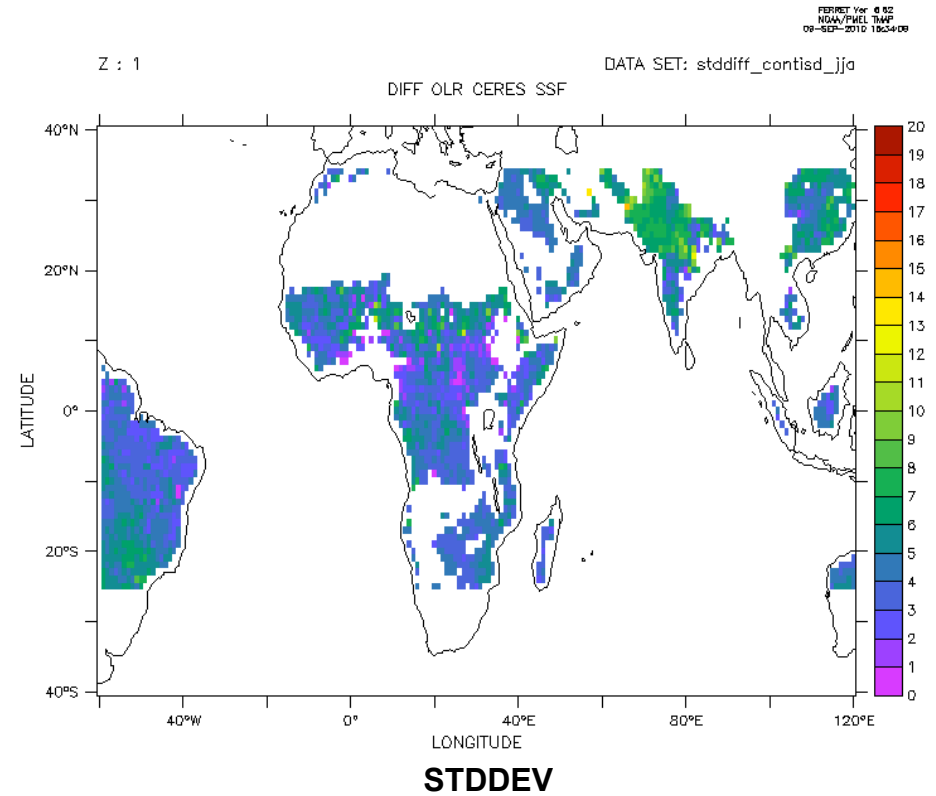
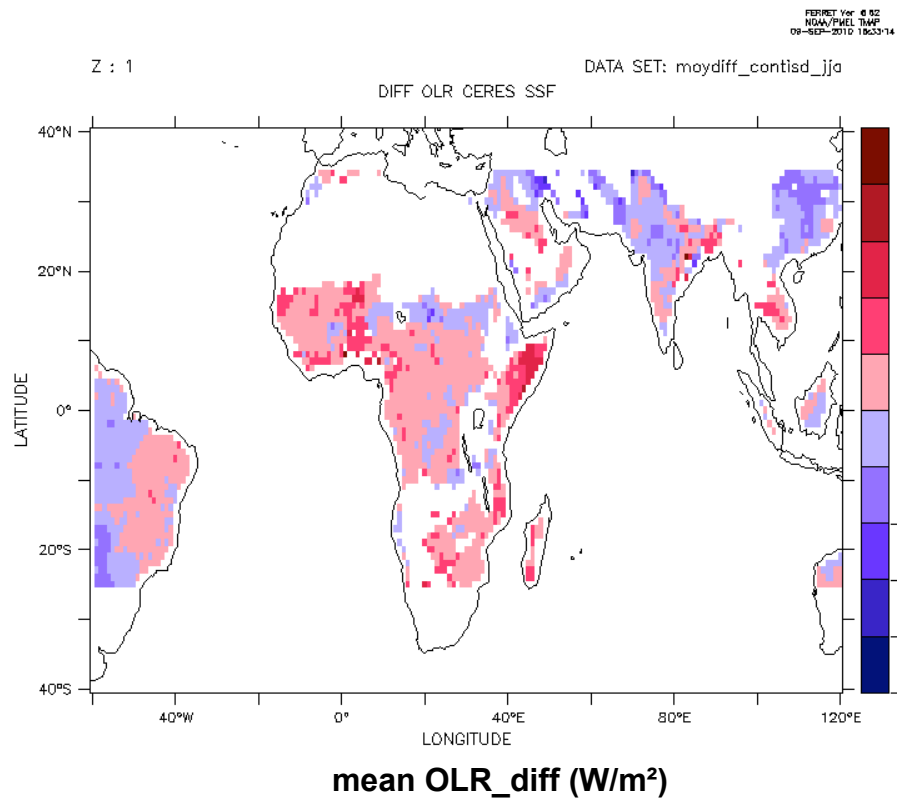
# Appendix 3

## Regression



# Appendix 4

## Regression



# Appendix 5

## Regression

